COMMUNITY EARTH SYSTEM MODEL (CESM)

Radiation

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Outline

- 1. Radiative Transfer in the Atmosphere
- 2. Methods for solving radiative transfer
- 3. How CAM solves radiative transfer: RRTMG
- 4. Sub-grid variability for clouds (McICA)
- 5. The CAM Aerosol-Radiation Interface
- 6. Radiative Forcing

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What does Radiation Do?

- Climate: energy balance of the planet
 - Gases: basic state & Impacts of greenhouse gases
 - Radiative transfer of heat between climate system components (mostly atmosphere and surface).
 - Clouds and condensed species (Aerosols) have unique challenges
- Weather: surface and atmospheric heating
 - What energy hits the surface
 - Clouds mediate this

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Spectrum of Solar Radiation (Earth)



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Water IR Absorption



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· JET Y

Cloud Radiative Effects in the Climate system

Two important properties of clouds: Clouds are White (SW), Clouds are Cold (LW)

Solar (cooling) & Infrared (heating)





Cloud Radiative Effects



IPCC 2013 (Boucher et al 2013) Fig 7.7

Aerosols

Clear sky in Beijing, July 2010, 17:00 LT looking west



Radiation Parameterization

Typical Parameterization:

- LW absorption/emission,
- SW scattering/absorption
- Cloud absorption and scattering (aerosols too).
- Plane-parallel. (3D much more expensive)
- Single-scattering approximation

Methods for Solving Radiative Transfer

Assumptions:

- Plane-parallel (single column)
- Single-scattering approximation

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Methods:

- Line-by-line
- Broadband
- Correlated-K



Methods: Line by Line

Typically 1000's of lines. Most accurate, but expensive (treat every line) Lines are partly 'empirical' (measured in lab, not theory)



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Attenuation Coefficient

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Methods: Broadband

Typically 10's of bands. Cost effective, but a big approximation Subject to significant errors



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Methods: Correlated-K

Typically ~100 bands. Accurate and cost effective Group bands by Absorption/Attenuation....



Attenuation Coefficient

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RRTMG Methods: Overview

- Correlated-k distribution for treating gases (sets spectral discretization)
- SW Transport calculations using two-stream (aka delta-eddington) in the single-scattering approximation in the SW.
- LW: same approach: neglect LW scattering
 - 2x faster without LW scattering,
 - LW scattering is about a 1-2% correction
 - LW is basically a 1-stream method computed at a single quadrature angle to represent full hemispheric irradiance.

Basics: Heating Rates CGILSS11

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Aside: Tropical Tropopause

Near tropopause: Transition between net radiative cooling (H2O) and warming (O3).

Zero Rad Heating = base of the Tropical Tropopause Layer

Where condensation heating is no longer balancing radiative cooling



Aside: Tropical Tropopause

GETTELMAN ET AL.: TROPICAL TROPOPAUSE LAYER RADIATION BALANCE



Can break down the heating rates by gases

• LW: H2O most important. Note CO2 rad heating in lower strat

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• SW: H2O (smaller), O3 in stratosphere, CO2 all the way through

Cloud Overlap

- RRTMG methods on sub-columns that include gases and condensed species (clouds and aerosols)
- Pre-computation of optical properties for clouds and aerosols,
 - Depends on some parameters (e.g. effective radius for cloud drops, with more complicated parameters for ice and aerosols)
- Sub-columns to treat sub-grid variability in clouds
 - Cloud Overlap
 - Correlation of water vapor and condensed cloud
- Monte Carlo Independent Column Approximation (McICA)
 - Sample the subc-olmns for efficiency

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Cloud Overlap



McICA

Monte carlo Independent Column Approximation

- Sub-grid variability of clouds (cloud fraction) requires radiation to "guess" cloud overlap.
- Assume clouds maximally overlap if they are in neighboring levels
- Assume clouds randomly overlap if there is a vertical gap between clouds
- Construct independent columns that satisfy the cloud fraction and cloud overlap model
- Sample the independent columns when performing the radiative transfer computation in each of the 128 correlated-k bands
- [Each independent column is applied to a different "color" of light]
- Enough monte-carlo samples converge to full radiative transfer though collection of independent columns.



Aerosols and Radiation

- Aerosol optical properties (like clouds) have slow variation across the spectrum
- No subgrid fraction
- Strongest net flux effect is in solar
 - Scattering and absorption
- Small net flux effect in the longwave
- But: significant stratospheric LW heating
 - Volcanic eruptions / geoengineering

Aerosol-Radiation Interface

• RRTMG

- Radiative transfer,
- McICA,
- gas phase optics.
- Condensed phase optics (in-cloud and aerosol optics) are 'parameterizations' based on microphysical state and are computed externally to RRTMG.
- Optics are assumed to be function of band, but constant for each correlated-k band inside of each band
- Optics are based on microphysical state of clouds/aerosols

Optics Based on Microphysics

Ice clouds

- Morphology fixed: shape recipe based on mid latitude cirrus (Mitchell)
- Size distribution

Liquid clouds

• Fully consistent mie calculations with mu and lambda parameters from cloud drop size distribution

MAM aerosol

- 3, 4, or 7-bins of internally mixed aerosol
- number-weighted average of index of refraction of composition
- Log-normal size distribution with fixed width and varying radius
- Water-vapor uptake

BAM aerosol

- Externally mixed with single-chemical composition
- Lognormal distribution with fixed (dry) radius and width
- Water-vapor update based on relative humidity

Liquid Cloud/Aerosol Optics





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Ice Optics



CAM uses a specified habit mix appropriate for mid-latitude cirrus (-20 to -35°C)

Radiative Forcing

"Change in net top-of-atmosphere flux due to change in composition/emissions/optics/surface"

- Instantaneous Radiative Forcing (IRF)
 - Net Flux change due to a change in composition given a specific time-sample-sequence of atmospheric states.
- Radiative Forcing (RF)
 - Same as IRF, except the stratosphere temperature is allowed to relax under the assumption of fixed stratospheric dynamics (but heating changes)
- Effective Radiative Forcing (ERF)
 - Allow land and atmosphere to respond, but fix ocean temperatures
- ERF w/TS correction
 - Like ERF, but recompute TOA flux with surface temperatures adjusted back to original case for the flux computation

How to Calculate RF

- Difference of 2 runs (but climate noise)
- Alternative method:
 - Additional 'diagnostic' radiation call
 - Can remove one gas at a time
 - Can also remove aerosols
 - Yields an RF, Difference of two of these is ERF
- Examples follow
 - Exercise will show how this 'rad_diag' works (namelist parameter)
- Also: PORT (Parallel Off-Line Rad Transfer)
 - Method to calculate column radiation off line for a collection of columns (like the whole planet)
 - Good for RF calculations

Volcanic ERF



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Schmidt et al 2018

Anthropogenic Aerosol ERF

A) Clearsky SW Aerosol Kernel

B) Allsky SW Aerosol Kernel



Aerosol Kernel of Wm-2/AOD from CAM5 Present - Pre Industrial Simulations To get ERF multiply by dAOD

C) Clearsky LW Aerosol Kernel

D) Allsky LW Aerosol Kernel



Summary

- CAM Solves Rad Transfer using Correlated-k (RRTMG)
- Condensed phase optics vary by species
 - Liquid, Ice, Aerosols
 - Optics (tau) fed into radiation (Not part of RRTMG)
- Sub-grid variability for clouds (McICA)
- CAM Has Capabilities to manipulate the radiation code, e.g. Radiative Forcing